Atmospheric Circulation



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Atmospheric Circulation

Key Topics

- Composition and Structure
- Solar Heating and Convection
- The Coriolis Effect
- Global Wind Patterns
- Storm Systems
- Climate Patterns
- Effects on Ocean



Atmospheric Circulation

<u>Key Concepts</u>

- \succ Earth's atmosphere is stratified and consists mostly of N₂ and O₂
- Differential heating of Earth's curved surface by incoming solar radiation produces an over-heated equator and under-heated poles
- Heating and cooling creates regions of low and high atmospheric pressure, respectively. Coupled with gravity, this drives atmospheric convection, i.e. global-scale, circulation system of winds
- The Earth has three major atmospheric circulation cells in each hemisphere – a total of six around the planet
- Earth's rotation causes moving air masses to curve left in the northern hemisphere and right in the southern hemisphere; this is called the Coriolis Effect
- Atmospheric circulation is responsible for the transfer of 2/3rds of Earth's surface heat from the equator to the poles
- > Spinning storm systems are of two types: Tropical and Extra-tropical
- Circulation of atmosphere and ocean moderates Earth's surface temperatures, and shapes weather and climate
- Surface winds and storms generate ocean currents and wind waves

Atmosphere Composition





Proportion of Gases in the Atmosphere

Key Ideas

Mostly consist of nitrogen (78%), oxygen (21%) and argon (1%)
Water vapor and carbon dioxide important minor components
Water vapor can be as high 4% by volume
Air has mass: 1 sq. foot column of the vertical atmosphere weighs 1 ton

Vertical Structure of Atmosphere Key Ideas Density-stratified air column 0 Most of air found in troposphere 0 Weather occurs in troposphere 0 Jet stream at top of troposphere Ozone found in stratosphere 0 Temperature inversions at the layer boundaries 0°C Exosphère 10000 Magnétosphère Altitude (km) 1000

#

250

Température (°K)

auche d'azone

300

900

lonosphère (thermosphère)

200

Mésosphere

Stratosphère

Troposphère

150

100



10 000 km

Annual Solar Energy Striking Earth



Incoming Solar Radiation at Top of Earth's Atmosphere



Annual Solar Radiation at Earth's Surface (kcal/cm2/year)

Key Idea:

 Global variation in the amount of solar energy striking Earth's surface is controlled by the *latitude*, *season*, *atmospheric conditions*, and *altitude*.



 Insolation is incoming solar radiation. The amount of insolation received at the surface of the earth is primarily controlled by the sun angle. Sun angle is a function of latitude and season.

Atmospheric Conditions Affect Amount of Sunlight Striking Earth's Surface

Key Idea

The amount of sun energy received at the surface of the earth can also be affected by cloud cover, dust, and other particulates that reflect and/or absorb incoming sunlight Sea surfaces are typically much more prone to cloud cover than land surfaces.





Uneven Solar Heating of Earth's Surface Causes Global-Scale Atmospheric Convection





Human-Scale Convection Process

Global-Scale Convection Process

Heat difference causes pressure differences in the overlying atmosphere
 Overheating of equatorial regions forms belt of low pressure
 Under-heating of polar regions creates centers of high pressure
 Pressure differences in lower atmosphere cause air masses to moves
 Air masses move from regions of high pressure to regions of low pressure



 Air masses move from regions of high pressure to regions of low pressure
 Severity of pressure gradient between adjacent regions of high and lows controls how strong of wind will blow between the high and low

Atmospheric Circulation Model of a Non-Spinning Earth



Key Ideas

- One cell per hemisphere
- Overheated equatorial air rises and moves horizontally aloft toward the poles
- Overcooled polar air sinks and moves horizontally at surface towards equator



Single-Cell Hemispheric Convection Model

Atmospheric Circulation Model of a Spinning Earth



Key Ideas Three cells per hemisphere Hadley, Ferrel, and Polar Similar convection process Smaller convective cells Trip Two surface convergence zones Two surface divergence zones



Triple-Cell Hemispheric Convection Model

- ✓ Spinning causes the Coriolis effect
- ✓ Coriolis effect deflects air currents



Objects deflect to the right in Northern Hemisphere
 Objects deflect to the left in Southern Hemisphere
 Moving air masses appear to have curved paths

Earth's Surface Winds





Land-free Circulation Model

- 1) Polar Easterly belts
- 2) Mid-latitude Westerly belts
- 3) Low-latitude Tradewind belts
- 4) Subpolar low and equatorial low pressure belts
- 5) Subtropical high pressure belt

Landmass Circulation Model

- 1) Landmasses break up the wind and pressure belts
- 2) High to mid-latitude pressure centers replace pressure belts
- 3) Equatorial low pressure belt
- 4) Seasonal shift of pressure centers



Convergence and Divergence Zones



Equatorial Convergence





High - Low Air Pressure Gradients, Wind, the Coriolis Effect, and Ground Friction

Pressure Gradient Force only (without Coriolis effect)

Pressure Gradient Force with Effect of Coriolis (in the upper atmosphere)

Pressure Gradient Force with **Effect of Coriolis and Ground Friction (in the lower atmosphere** at/near ground surface) (a) Pressure Gradient Force only:



(b) Upper Atmosphere—Pressure Gradient Force and Coriolis Effect:



(c) Lower Atmosphere—Pressure Gradient Force, Coriolis Effect, and Friction:



Earth's Climate Zones Precipitation (P) evaporation (E) Latitudinal Climate Zones: Fronts Subtropics **Regions dominated by low** pressure (Fronts and ITCZ Hadley ITCZ cells =favors precipitation) Subtropics 30° Regions dominated by high Fronts pressure (Subtropics and poles = favors evaporation) 1) Polar high pressure = cold and dry 2) Sub-polar low pressure = polar fronts = cool and wet 3) Sub-tropical high pressure = horse latitudes = warm and dry 4) Equatorial low pressure = ITCZ or doldrums = hot and wet

Divergence Versus Convergence Air Flow at High and Low Pressure Centers



Atmosphere pressure differences cause air masses to moves = Wind
 Air masses move from regions of high pressure to regions of low pressure
 High pressure centers have rotating downdrafts of drying air
 Low pressure centers have rotating updrafts of moistening air

Adiabatic Effects of Vertical Air Mass Movement



Temperature and Volume Changes as a Function of Vertical Air Mass Movement <u>= adiabatic</u> cooling (uplift) or warming (downdraft) Key Ideas



Water Vapor Saturation Level as a Function of Changing Temperature of Air Mass

Ascending (rising) air expands, cools, and becomes moister
 Descending (falling) air contracts, heats, and becomes dryer
 Water vapor in rising and cooling air will condense into clouds
 Further rising and cooling of cloud-rich air will lead to precipitation

Atmospheric Humidity versus Temp



Maximum Vapor Pressure (Absolute Humidity Limit) for Air Masses of Different Temperatures



Relative Humidity for Air Masses of Different Temperatures with a Given Absolute Amount of Water Vapor

Key Ideas Ascending (rising) air expands, cools, and becomes moister Descending (falling) air contracts, heats, and becomes dryer Water vapor in rising and cooling air will condense into clouds Further rising and cooling of cloud-rich air will lead to precipitation

Weather at Divergence and Convergence Zones

Polar Divergence

High evaporation Variable winds Cold, harsh, dry weather Subpolar Convergence **Heavy precipitation** Winter storm fronts Stormy, wet, cool weather Subtropical Divergence **High evaporation** Variable winds and Calms ✓ Warm, mild, dry weather **Tropical Convergence** ✓ Heavy precipitation **Light winds and Calms Tropical cyclone nursery** Stormy, wet, hot weather



The Jet Stream and Pressure Centers



Key Ideas

 Narrow, fast-moving, ribbons of high-level wind moving west to east
 Found between atmospheric cells
 Controls position and movement of stormy low pressure systems (above troughs) and fair weather high pressure systems (below ridges)







Daily reversal of local coastal winds

- Coastal breeze direction controlled by local differences in atmospheric pressure between air masses over land versus sea
- Nighttime offshore wind (Land = High; Sea = Low

Daytime onshore wind (Land = Low; Sea = High





Seasons, Surface Winds and Weather



Key Idea

Seasonal changes in Earth's axis in respect to the sun cause latitudinal migrational shifts in several atmospheric elements:

- Pressure centers
- Wind belts
- Jet streams
- Intertropical Convergence Zone (ITCZ)
- Large-scale weather patterns

Pressure Systems and Wind Patterns Averaged January Pattern





Wind Speed (meters/sec

key Ideas

The ITCZ is shifted to its maximum southward position
Polar lows dominate the northern Pacific and Atlantic Oceans
Subtropical highs dominate south Pacific and Atlantic Oceans
Winter in the Northern Hemisphere; summer in the S. Hemi.

Pressure Systems and Wind Patterns Averaged July Pattern





Key Ideas

The ITCZ is shifted to its maximum northward position
Subtropical highs sit over the north Pacific and Atlantic Oceans
Southern Ocean westward wind belt at maximum strength
Summer in the Northern Hemisphere; Winter in the S. Hemi.



Local atmospheric conditions that prevail over a period of days to weeks



Spinning Air Mass Disturbances

 Tropical Cyclones
 Extratropical Cyclones

Solar Energy Powers the Both the Atmospheric and Hydrologic Cycles



Solar Energy Causes Evaporation of the Ocean Surface Waters
 ✓ 1 meter of ocean surface is evaporated each year!
 ✓ Most precipitation falls back into the ocean
 ✓ Precipitation over land plays huge part in weathering and erosion



Several Distinctive Regional Air Masses found Across the Hemisphere
✓ Each air mass gets its characteristics from the surface beneath it
✓ Differences in surface type (marine versus continental)
✓ Differences in surface temperature (equator vs. subtropical vs. polar)
✓ Each air mass has different density; they do not mix when they collide.
✓ Stormy weather is the result of rising and colliding air masses.



Tropical Cyclones

Tropical Cyclones

- 1) Tropical Cyclones are known as hurricanes in the Atlantic Ocean, typhoons in the Pacific Ocean and cyclones in the Indian Ocean. Very extensive, powerful, and destructive type of storm. 3) This type of storm develops over oceans 8° to 15° North and South of the equator.
- 4) Hurricanes draw their energy from the warm water of the tropics and <u>latent heat</u> of <u>condensation</u>.



Necessary Conditions for Cyclone Development:

- 1) Must originate over ocean water that is least 26.5 °C.
 - Hurricanes feed off the <u>latent heat of water</u> hotter the better!
- 2) Have an atmosphere that cools quickly with height.
 - This creates potentially unstable conditions that builds storms.
- 3) Low vertical wind shear.
 - \checkmark Winds at all levels of the atmosphere from the ocean up to 30,000 feet or higher are blowing at the same speed and from the same direction.
- 4) No closer than 500 kilometers to the equator.
 - The Coriolis Force is too weak close to the equator.
 - ✓ It is the Coriolis Force that initially makes the cyclone spiral and maintains the low pressure of the disturbance.
- 5) An upper atmosphere high pressure area above the growing storm.
 - ✓ The air in such high pressure areas is flowing outward. This pushes away the air that is rising in the storm, which encourages even more air to rise from the low levels.
- 6) Hurricanes will not always form in these conditions. However, a will hurricane only form if these conditions are present.

Anatomy and Behavior of a Hurricane



Warm, humid surface winds spiral towards eye.
 Strongest winds occur in the eye wall at the surface.
 Air in the eye sinks which inhibits wind and cloud formation
 Body of hurricane divided into concentric rain bands
 Surface rotation direction depends on hemisphere
 All hurricanes move toward the west

Life Cycle of Tropical Cyclones

Formation

Tropical Disturbance to Depression Weak to moderate winds Little to no rotation

2) **Prematurity**

Tropical Storm Moderate to strong winds Moderate rotation

3) Full Maturity



✓ Hurricane Very Strong winds Rapid rotation with eye

4) Decay



 Entire cycle typically lasts between 1 to 2 weeks





Global Tropical Cyclone Tracks



Which ocean basin has the most tracks? Why? Which ocean basin has the least tracks? Why?

Hurricane Intensity Scale











Hurricane Rita was the strongest hurricane ever recorded in the Atlantic Basin





www.windyty.com

Mid-Latitude Cyclonic Systems

What is a Mid-Latitude Cyclone?

 ✓ A high- to mid-latitude cyclonic low pressure system

✓ Typically form from fall to spring

✓ IT IS NOT A HURRICANE OR TROPICAL STORM

✓ Forms from the swirling convergence of a cold and warm air mass along the polar front

 ✓ Associated with high winds, clouds and precipitation

✓ Typical size of mid-latitude
 cyclone = 1500-5000km in diameter





Development of Mid-Latitude Storm Systems



1) Frontal Wave Develops



2) Instability "Kink" Forms







3) Fully-developed Cyclone

Example

Cold and Warm Fronts



Spot the Three Frontal Systems -



Where are the centers of systems?
 Centers of low pressure?
 Sense of rotation?
 Regions of warm and cold air masses?



Development of Ocean Currents





Prevailing Surface Winds





Wind-forced Ocean Currents



Resultant Ocean Gyres

Mind – Maye Connection



Surface Wind Speeds

Highest = Westerlies
Lowest = Doldrums/ ITCZ

Wave Heights
> Highest = High Latitude
> Lowest = Low Latitudes

KEY POINTS

Ocean seas and swell are direct result of ocean surface winds
 The stronger the wind, the larger the seas and swell

Development of Ocean Wind Waves



Prevailing Surface Winds



NOAA/NCEP Feb 24 2010 12z 00 hr fcst

0 1 2 3 4 5 6 7 8 9 101214161820253035405075 Wave Watch III Sig. Wave Height(ft) and Direction

GMT Feb 24 13:55:48 2010 La Jolla Sunting - http://lajoilasuri.org/

Resultant Ocean Wind Swell

Atmospheric Circulation Review of Key Concepts

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- The Earth has three major atmospheric circulation cells in each hemisphere – a total of six around the planet
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- Atmospheric circulation is responsible for the transfer of 2/3rds of Earth's surface heat from the equator to the poles
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Discussion



Preparation for Next Lecture – Ocean Circulation

- 1) Read Chapter on Ocean Circulation and Currents
- 2) Review End-of-Chapter Questions and Exercises
- Review Instructor's classroom website for:
 - ✓ Course Schedule
 - Lecture Notes
 - ✓ Lecture Presentation



